

Observations of enhanced nonlinear instability in the surface reflection of internal tides

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Received 4 February 2013; revised 4 March 2013; accepted 5 March 2013; published 29 April 2013.

[1] Enhanced vertically standing waves formed by the superposition of two upward and downward going near-diurnal (D_1) waves are observed during one semidiurnal (D_2) spring tide in an approximately 75 day long velocity record from the northeastern South China Sea. Bicoherence estimates suggest that the enhanced D_1 waves are likely due to nonlinear parametric subharmonic instability of D_2 internal tides. The timescale for energy growth by an order of magnitude is about 2.5 days for these waves. In addition to subharmonics, higher harmonics D_4 ($=D_2+D_2$) and a mean flow are generated by a different nonlinear interaction during the same D_2 spring tide. The separation of coherent from incoherent internal tidal signals and a rotary spectral decomposition in the vertical direction reveal that D_2 waves with opposite vertical propagation directions in the region of internal tide reflection from the surface may be responsible for the pronounced nonlinear instability. **Citation:** Xie, X., X. Shang, H. van Haren, and G. Chen (2013), Observations of enhanced nonlinear instability in the surface reflection of internal tides, *Geophys. Res. Lett.*, **40**, 1580–1586, doi:10.1002/grl.50322.

1. Introduction

[2] It is well known that semidiurnal internal tides (IT _{D_2} , internal waves with semidiurnal frequency D_2) dominate the internal wave spectrum in the ocean. Many observations and numerical models have suggested that IT _{D_2} can transfer their energy to a series of higher harmonics (e.g., D_4 , D_6) via nonlinear (superharmonic) instability [e.g., van Haren *et al.*, 2002; Lamb, 2004; Xie *et al.*, 2008]. On the other hand, they can also lose energy to waves with frequencies near the diurnal frequency D_1 ($=D_2/2$) and higher vertical wave number k_z through parametric subharmonic instability (PSI) [e.g., Hibiya *et al.*, 2002; van Haren, 2005; Alford *et al.*, 2007]. Although in theory the classic PSI is thought to be a weakly nonlinear process [McComas and Bretherton, 1977], several recent numerical models [Gerkema *et al.*, 2006; Korobov and Lamb, 2008] and field observations [Carter and Gregg, 2006; Xie *et al.*, 2011] suggest that PSI of IT in the ocean may turn into a rapid, strong instability.

[3] Two regions where pronounced nonlinear instability may occur are suggested by numerical models and laboratory

experiments: one is the boundary where internal wave beams are reflected [Javam *et al.*, 2000; Mercier *et al.*, 2012] and the other is the region where two beams intersect [Teoh *et al.*, 1997; Korobov and Lamb, 2008]. In the ocean, few observations exist on nonlinear effects in the above mentioned two regions. Although IT can be modeled by a purely horizontally propagating (vertical) mode in a flat-bottomed ocean [Leblond and Mysak, 1978], this model is destroyed when the water depth varies. Then, one of the downward and upward propagations will be dominant. So in a real ocean with varying water depth, IT will always be a vertical traveling wave. In this paper, moored current measurements in the northeastern South China Sea (SCS) are used to investigate nonlinear effects between two oppositely propagating internal tidal waves in the vertical direction in the near-surface reflection area of IT. Our observations suggest that PSI from IT _{D_2} can be rapidly enhanced in the reflection region. A mean flow caused by the nonlinear interaction between these two waves is also observed.

2. Data and Methods

[4] Ocean currents are evaluated from one mooring deployed at 20.58°N, 118.41°E, where the water depth is 2474 m. The mooring consisted of an upward looking 75 kHz acoustic Doppler current profiler (ADCP) at ~450 m (data from 436 to 52 m, 16 m bins) and five Aanderaa current meters at 800, 1000, 1500, 2000, and 2300 m. In this paper, the ADCP data and current meter data from 2000 and 2300 m are used. The mooring period was from August 2000 to March 2001. Unfortunately, the ADCP worked only during the first 75 days and current meters at 800, 1000, and 1500 m gave false data due to some incorrect settings before deployment. The sampling rates of ADCP and current meters were once per 15 and 30 min, respectively. Data of the two current meters at 2000 and 2300 m are linearly interpolated onto the 15 min sample interval to match those of the ADCP data.

[5] In the SCS, diurnal (D_1) internal tides (IT _{D_1}) are often dominant in the internal wave spectrum [Lien *et al.*, 2005; Xie *et al.*, 2009]. For a 75 day long record, one can easily separate inertial (frequency: $f=2\pi/(34\text{ h})$), D_1 , and D_2 motions. Using a second-order Butterworth filter, all current data are band-pass-filtered to extract the D_1 and D_2 components, using filter bounds $[0.9O_1, 1.1K_1]$ and $[0.86M_2, 1.14M_2]$ (O_1 , K_1 , and M_2 are the primary tidal frequencies in the D_1 and D_2 bands), respectively. The band-passed D_1 and D_2 motions are further decomposed into coherent and incoherent components by applying a sharply harmonic filter [van Haren, 2004]. The coherent component containing barotropic tides and coherent IT (hereafter referred to as CIT) can be obtained by

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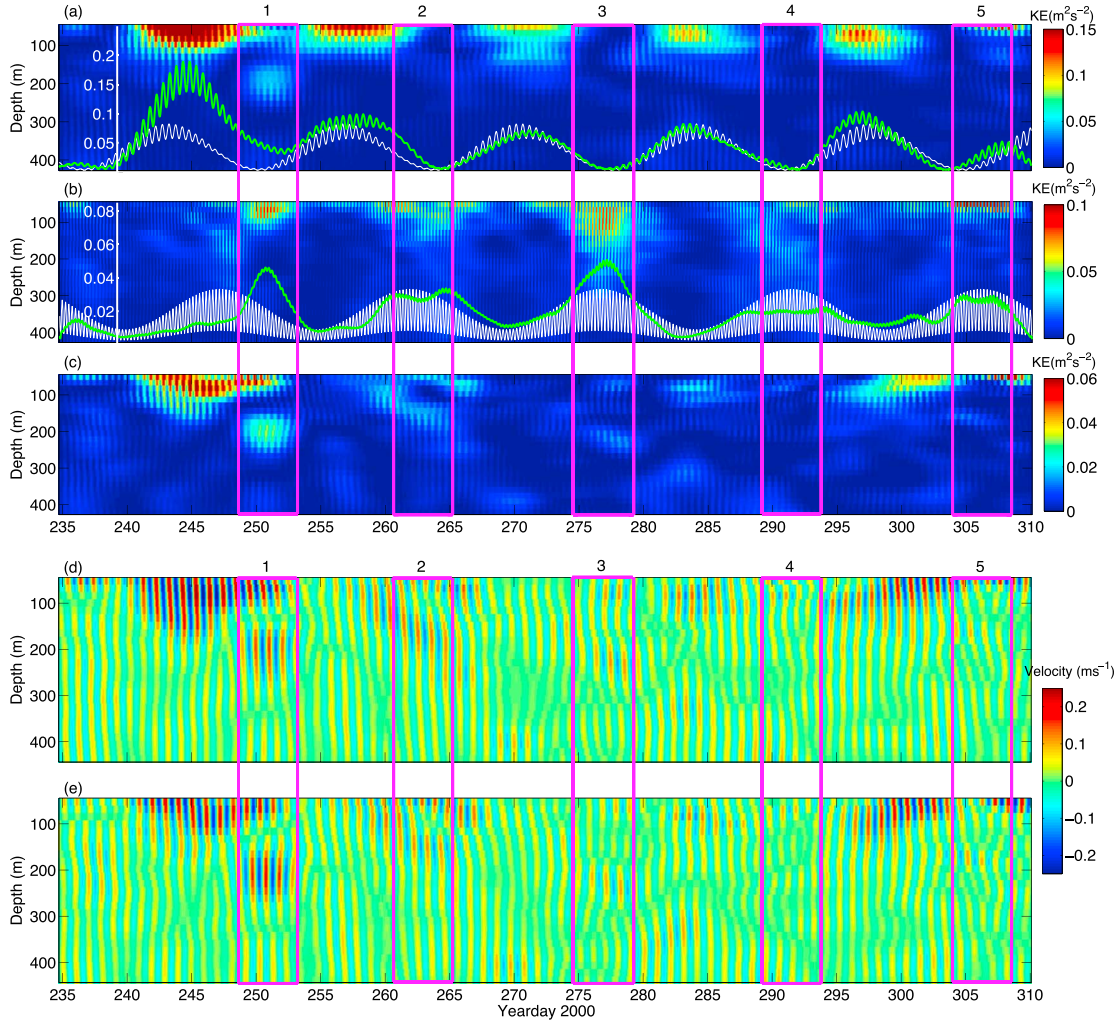


Figure 1. Time series of kinetic energy (KE) and velocity from the ADCP data. (a) KE_{D1} . (b) KE_{D2} . (c) KE of diurnal ICIT (KE_{icD1}). (d) Zonal and (e) meridional velocities of $ICIT_{D1}$. The green and white curves in Figure 1a are KE_{D1} and KE_{O1+K1} averaged above 120 m, respectively. The green and white curves in Figure 1b are KE_{D2} and KE_{M2+S2} averaged above 120 m, respectively. The white vertical scale in Figures 1a and 1b indicates the magnitude of the green and white curves. In all panels, the purple rectangles indicate the periods of SSTs. The SSTs are numbered. Note the different color scales in Figures 1a, 1b, and 1c.

$$u_c = \sum_n U_n \cos(\omega_n t + \varphi_n)$$

$$\omega_n = Q_1, O_1, K_1 (= 2N_2, M_2, S_2 \text{ for } D_2 \text{ band}), \quad (1)$$

where U_n is the amplitude of the n th constituent and φ_n is its phase, similar for meridional component v_c [van Haren, 2004]. The incoherent IT (hereafter referred to as ICIT) component is then defined as

$$u_{ic} = u_{\text{Bandpass}} - u_c, \quad (2)$$

similar for the v component. In equation (1), six frequencies are chosen, which all have a signal-to-noise ratio larger than 1.5.

3. Results

3.1. Diurnal ICIT

[6] Figures 1a and 1b show the time series of D_1 and D_2 kinetic energy, actually current variance, KE_{D1} and KE_{D2} ($KE = u^2 + v^2$), respectively. KE_{D1} and KE_{D2} values in the

upper 120 m show approximate fortnightly spring-neap cycles. The diurnal 14 day cycle is induced by the superposition of O_1 and K_1 (Figure 1a), while the semidiurnal 14 day cycle is mainly induced by the superposition of M_2 and S_2 (Figure 1b). However, there is a significant variation in semidiurnal spring tides (SSTs). Especially, the largest KE_{D2} around the first SST (SST-1) appear on year day 250, having a significant shift with respect to the peak of coherent $M_2 + S_2$ signals. Hereafter, SSTs (marked by rectangles in Figure 1) refer to the periods with the largest KE_{D2} averaged over all observation depths above 120 m around the times of spring tide induced by coherent $M_2 + S_2$ signals. The above shift is partly due to the modulation of other coherent constituents such as $2N_2$. On the other hand, the shift may also be due to the varying background condition, which can generate incoherent components to destroy the 14 day spring-neap cycle [van Haren, 2004]. This is also likely to be one of the reasons for the generation of strong $ICIT_{D1}$ around year day 245 (Figure 1c). Noticeably, a relatively weak growth of KE_{D1} occurs near 200 m on year days 249

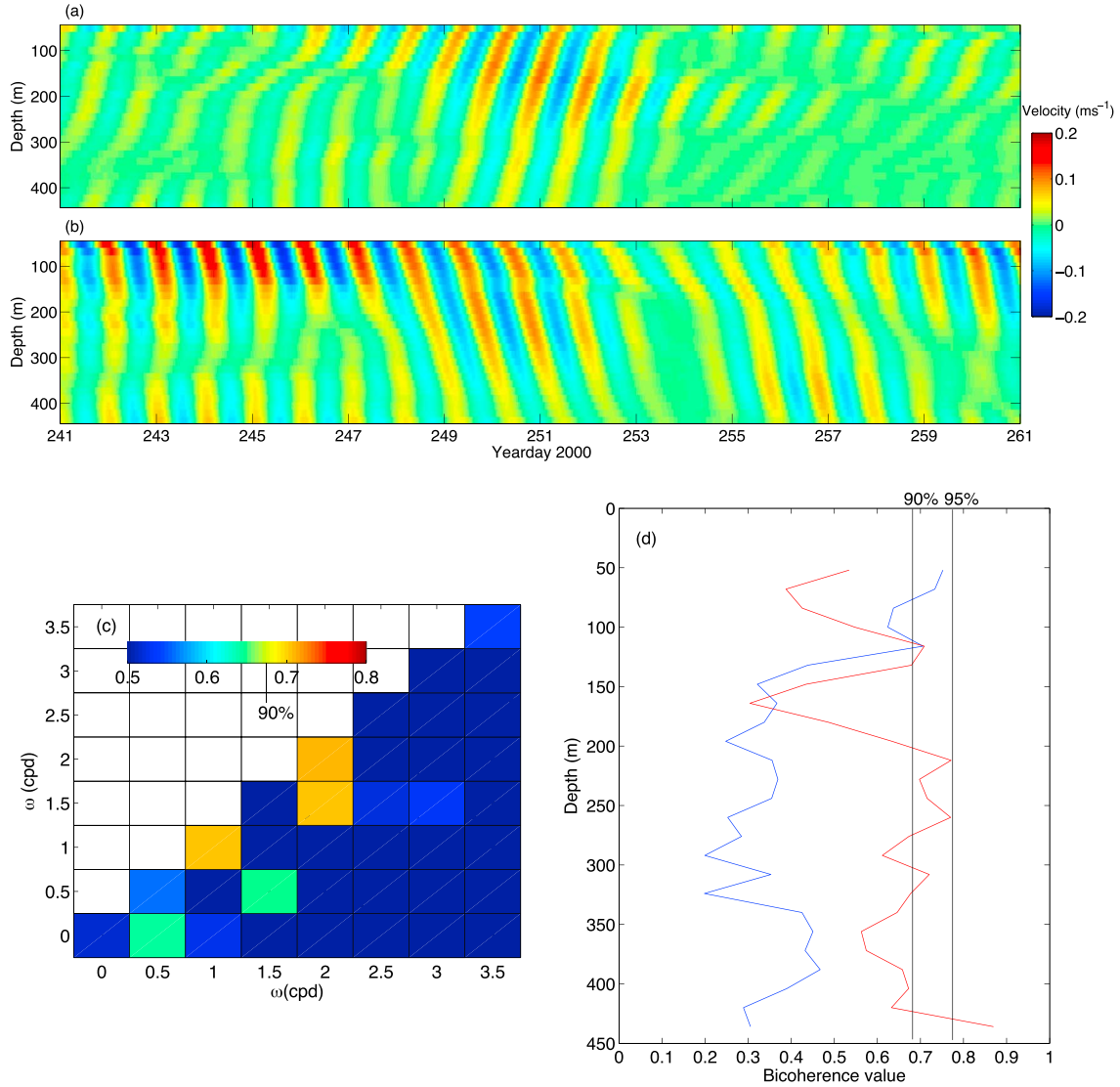


Figure 2. (a) CW and (b) ACW motions of ICIT_{D1} from the ADCP data on yeardays 241 to 260. (c) Average bicoherence of u and v at 116 m during SST-1. Note that all values less than 0.5 are set to 0.5 for clarity. (d) Bicoherence values around [1, 1] cpd (blue) and [2, 2] cpd (red) are plotted as a function of depth. The vertical lines give 90% and 95% significance levels.

to 252 corresponding to diurnal neap tide (Figures 1a and 1b). After removing the coherent $O_1 + K_1 + Q_1$ component, the D_1 14 day cycle disappears (Figure 1c). Enhanced ICIT_{D1} only occur during some SSTs. Specifically during SST-1, ICIT_{D1} near 200 m and D_2 motions above 120 m are synchronously enhanced and weakened (Figure 1b). In both the zonal and meridional incoherent velocity fields of D_1 (Figures 1d and 1e), the enhanced D_1 motions show large k_z ($\approx 2\pi/300$ m). Furthermore, a checkerboard (vertically standing) pattern is clearly identified on yeardays 249 to 252, which indicates that the signals are the sum of upward and downward propagating waves. During other SSTs, signals with high k_z ($> 2\pi/300$ m) are also observed, but no clear checkerboard pattern is found. In contrast, during D_1 springs ICIT_{D1} show relatively low k_z ($< 2\pi/300$ m).

[7] In order to confirm the bidirectional propagation of ICIT_{D1} during SST-1, incoherent velocity of D_1 is decomposed into clockwise (CW) and anticlockwise (ACW) rotary motions with depth, which represent upward and downward phase propagation, respectively [Leaman

and Sanford, 1975]. The decomposed velocity fields are shown in Figure 2. As expected, both upward (CW) and downward (ACW) traveling signals with comparable amplitudes are clearly identified on yeardays 249 to 252 (Figures 2a and 2b). In contrast, ICIT_{D1} during diurnal springs (yeardays 241–246 and 255–260) are dominated by downward propagation.

3.2. Bicoherence Spectra

[8] In section 3.1, we show that the features of ICIT_{D1} observed during SST-1 agree well with those of subharmonics as expected from PSI theory [McComas and Bretherton, 1977]. In order to provide evidence of near-diurnal subharmonic energy transfer, bicoherence (normalized bispectrum) is computed following Carter and Gregg [2006]. Because CIT_{D1} may contaminate this spectral estimation, coherent $O_1 + K_1 + Q_1$ signals are removed from the raw velocity record before computing bicoherence. Here, only a 6 day long (yeardays 247–252) velocity record around SST-1 is used for the estimate. The 6 day long time series is divided into five segments

with 50% overlap. Each 2 day time series segment is smoothed by a Kaiser-Bessel window, yielding a spectral estimation with ~ 10 degrees of freedom [Emery and Thomson, 2001].

[9] Figures 2c and 2d show averaged bicoherence of u and v velocity components. The bicoherence value at a particular frequency pair $[\omega_1, \omega_2]$ represents the strength of the triad wave-wave interaction at frequencies ω_1 , ω_2 , and $\omega_1 + \omega_2$. Analysis at 116 m shows significant bicoherence around the frequency pair $[1, 1]$ cpd (Figure 2c), which is evidence for near-diurnal PSI energy transfer of D_2 waves. Significant bicoherence also appears around $[2, 2]$ cpd, suggesting that quarter-diurnal (D_4) waves, which are also enhanced during SST-1 (see Figure 3g), may be generated by self-nonlinear interaction of D_2 waves, namely $D_2 + D_2 = D_4$. For subharmonic instability of D_2 waves, significant bicoherence results at the 90% level only appear above 120 m (Figure 2d). This is consistent with the depth range with the largest KE_{D2} (see Figure 1b). However, large bicoherence values around $[2, 2]$ cpd mostly appear below 200 m (Figure 2d).

3.3. Bidirectional Propagation of D_2 Waves

[10] In sections 3.1 and 3.2, it was shown that $ICIT_{D1}$ with a checkerboard pattern are likely generated by PSI of D_2 waves. Although we cannot confirm that other enhanced $ICIT_{D1}$, such as the weaker signals observed during SST-2, are also excited via PSI of D_2 waves, these signals are much weaker than those observed during SST-1. To explore the possible reason, primary waves (D_2) are investigated. Since IT mainly propagate westward in the northeastern SCS [e.g., Lien *et al.*, 2005], zonal velocities of D_2 around three SSTs (SST-1, SST-2, and SST-5) are presented. Velocity signals at different SSTs reveal large differences (Figures 3a and 4a). During SST-1 and SST-2, no clear phase propagation is observed above 400 m, while signals during SST-5 show marked upward phase propagation (Figure 4a). During SST-2, velocities at small (50–400 m) and large (2000 and 2300 m) depths are often out of phase, implying the dominance of local vertical mode-1. Although out-of-phase shallow-deep flows also appear during SST-1 and SST-5, higher-mode D_2 waves can be clearly identified. The higher-mode D_2 waves are primarily present in the form of ICIT because of their significant phase variation in the vertical direction (Figures 3c and 4c).

[11] Around SST-1, the phase of $ICIT_{D2} + CIT_{2N2}$ indicates downward phase propagation. On the contrary, CIT_{M2+S2} are dominated by waves with low k_z , showing upward phase propagation (Figure 3b). As coherent $2N_2$ signals are weak, $ICIT_{D2} + CIT_{2N2}$ is hereafter referred as $ICIT_{D2}$. At the same SST, the bidirectional vertical propagation of IT_{D2} can also be clearly identified when D_2 currents are decomposed into CW and ACW rotary motions (Figures 3d and 3e). $ICIT_{D2}$ with significant downward phase propagation are not found during SST-2 and SST-5 (Figure 4c), neither during SST-3 and SST-4 (not shown). Since during these SSTs CIT_{M2+S2} with larger amplitude propagate upward (Figures 4b and 4c), CW rotary motions often dominate over ACW rotary motions (Figures 4d and 4e). Furthermore, few events are observed with significant downward phase propagation in ACW signals during these periods. Thus, during SST-1, strong subharmonics may be generated via interaction between two vertically oppositely traveling D_2 waves. Their rapid energy growth is associated with the amplitude and phase of the two D_2 waves. This is

because on one hand during other SSTs, near-diurnal subharmonics are not effectively enhanced due to weak amplitude or absence of downward (phase) propagating waves. On the other hand, the bidirectional propagation of D_2 waves also occurs on yeardays 246 to 248 (Figures 3b, 3c, 3d, and 3e), during which upward propagating CIT_{M2+S2} and downward propagating $ICIT_{D2}$ above 120 m are nearly out of phase, but there are no enhanced subharmonics. In contrast, when strong subharmonics occur on yeardays 249 to 252, two waves are approximately in phase above 120 m where a strong interaction at $D_1 + D_1 = D_2$ is likely to occur (see Figure 2d).

[12] For the enhanced near-diurnal subharmonics on yeardays 249 to 252, the timescale of their energy growth $KE/(dKE/dt)$, where subharmonic KE is averaged over half a vertical wavelength (150–300 m; see Figures 1d and 1e), is estimated to be about 2.5 days (Figure 3f). This means that the energy increases by an order of magnitude in about 2.5 days. This characteristic timescale is roughly consistent with that predicted by the numerical model of Gerkema *et al.* [2006] and that found in the field observations of MacKinnon *et al.* [2013] near the critical latitude of 28.8° where the most effective PSI process is expected.

3.4. Higher Harmonic and Mean Flow

[13] In section 3.3, it was shown that rapidly growing near-diurnal subharmonics are associated with the bidirectional (vertical) propagation of IT_{D2} . Bidirectional propagation of IT near the surface may be due to the surface reflection of IT. Using small-amplitude expansion, Tabaei *et al.* [2005] theoretically discussed nonlinear effects in internal wave reflection from a boundary. They show that nonlinear interaction becomes pronounced in the reflection region where two IT beams meet. Thereby, a series of higher harmonics and a mean flow are generated. Recent laboratory experiments by Mercier *et al.* [2012] verify the generation of higher harmonics and a mean flow at the location where an IT beam reflects from the surface. To investigate whether these two motions are also generated when two oppositely traveling D_2 waves interact, the first higher harmonic (D_4) and subinertial velocities are extracted from the present data. As expected, D_4 waves are enhanced around SST-1 (Figure 3g). The starting time of the increase is around yearday 246, when the vertical propagation direction of $ICIT_{D2}$ is opposite to that of CIT_{M2+S2} (Figures 3b and 3c). Meanwhile, a westward (negative) mean flow restricted to the upper layer emerges from the background eastward flow (Figure 3h). Since IT in our observation area are expected to travel westward [e.g., Lien *et al.*, 2005], the direction of wave propagation is the same as that of the mean flow.

[14] The westward mean flow and elevation of D_4 waves, along with strong subharmonics (Figure 3f), disappear around yearday 252. Interestingly, the depth ranges where KE_{D4} is elevated are different in different periods (Figure 3g). On yeardays 246 to 248, D_4 waves are enhanced only above 100 m, where D_2 velocity is small (Figure 3a). After yearday 248, the near-surface elevation of KE_{D4} disappears. Instead, D_4 waves are enhanced below 180 m. The elevated KE_{D4} below 180 m occurs nearly synchronously with the enhanced subharmonics and near-surface elevation of KE_{D2} (Figure 3f). These significant variations of D_4 waves may be due to different nonlinear processes. On yeardays 246 to 248, the

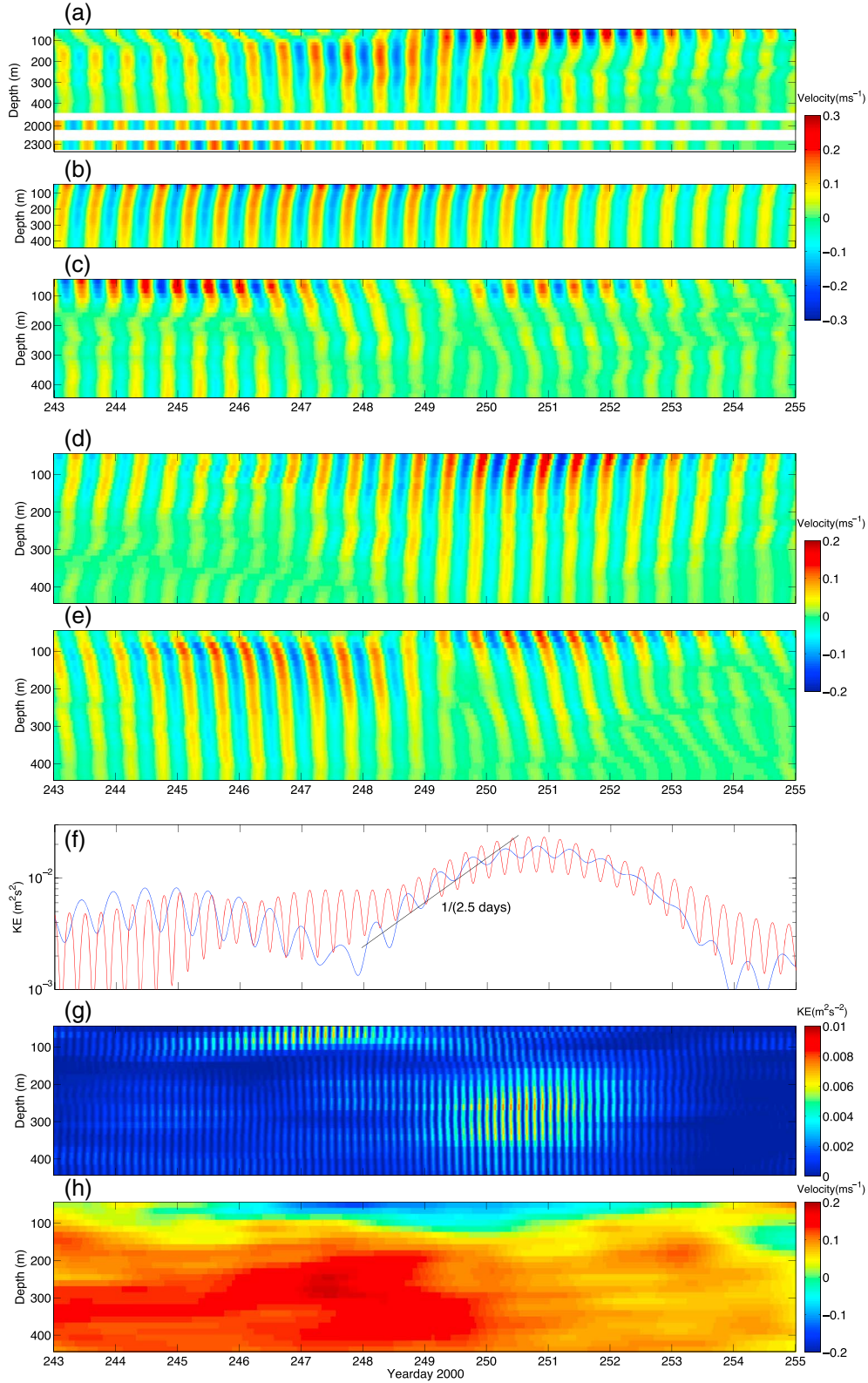


Figure 3. Time series of KE and velocity on yeardays 241 to 260. (a) Band-pass-filtered D_2 velocity. Values at 2000 and 2300 m are multiplied by a factor of 3. (b) Coherent $M_2 + S_2$ and (c) $ICIT_{D_2} + CIT_{2N_2}$ for the ADCP data only. (d) CW and (e) ACW motions of total D_2 signals from the ADCP data. (f) $KE_{D_2}/3$ averaged over the depth bins above 120 m (red line) and KE_{icD_1} averaged over the depth range of 150–300 m (blue line). The oblique line indicates the subharmonic energy growth rate with $1/(2.5 \text{ days})$. (g) KE_{D_4} . (h) The low-pass-filtered mean flow using the Butterworth filter with a cutoff frequency of $1/(3 \text{ days})$. Data in Figures 3a, 3b, 3c, and 3h are computed for the zonal velocity.

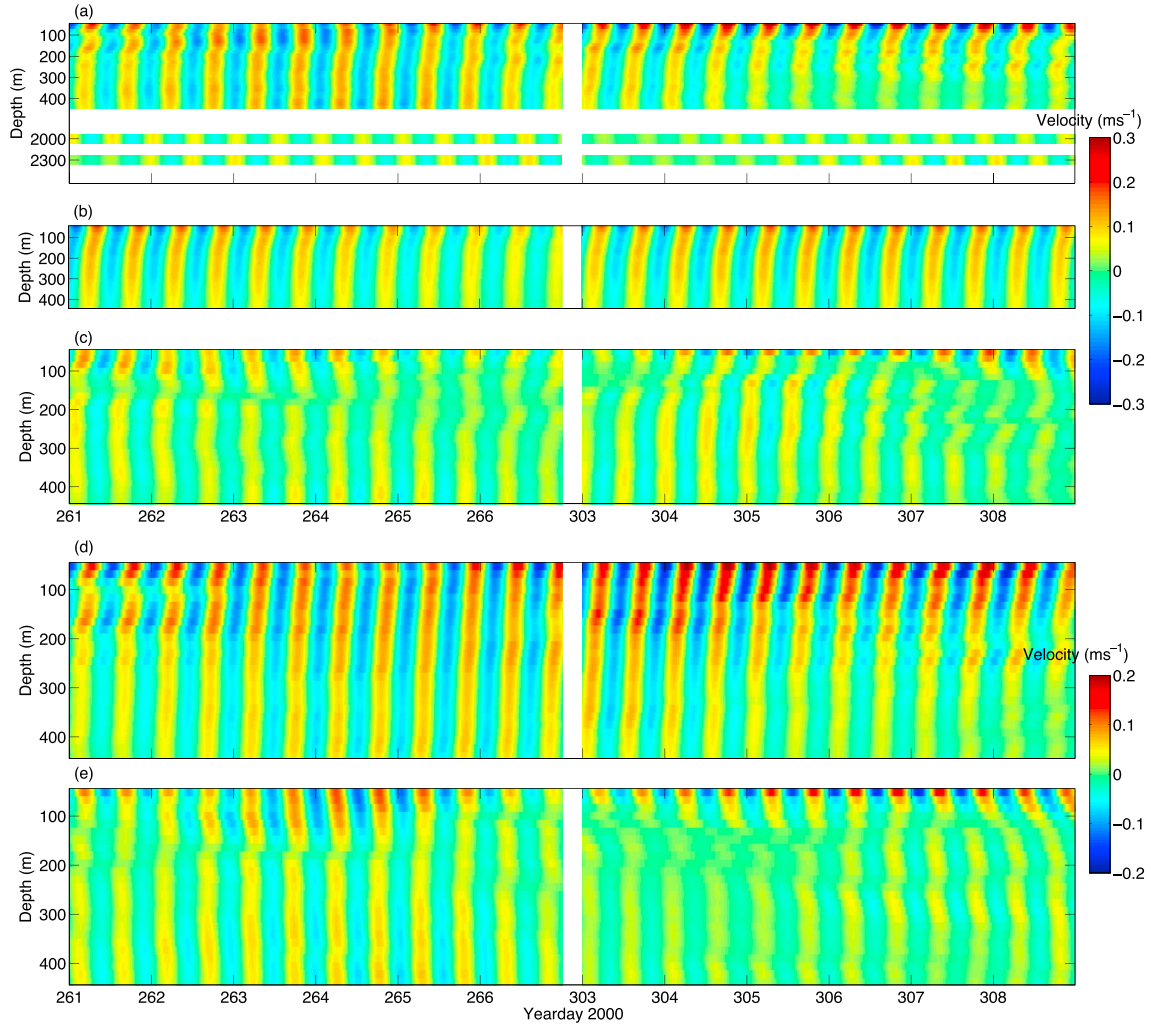


Figure 4. Similar to Figures 3a, 3b, 3c, 3d, and 3e but for the periods around SST-2 and SST-5.

generation of D_4 waves and the westward mean flow may be due to weakly nonlinear wave-wave interaction when the primary waves (D_2) are relatively weak, as shown by *Tabaei et al.* [2005]. After yearday 248, strong subharmonics and D_4 waves are likely the result from strong nonlinear instability when the primary waves are enhanced near the surface. However, it is unclear why the depth range with the elevated D_4 waves is out of the range of near-surface intensification of D_2 waves on yeardays 249 to 252.

[15] Two features of the mean flow, namely the confinement of the flow to the upper layer and the consistence of internal tidal horizontal propagation and the flow directions, agree with the laboratory results by *Mercier et al.* [2012]. So it seems that the surface reflection of IT_{D_2} may occur near our observation location, thereby generating strong subharmonics caused by PSI and a westward mean current. We note that *Mercier et al.* [2012] did not show the occurrence of PSI. This is probably because their experiments are without rotation, which plays an important role in improving subharmonic energy growth [MacKinnon and Winters, 2005].

4. Summary

[16] Current profiles have been presented suggesting the occurrence of pronounced nonlinear interactions in the

region where two oppositely (upward and downward) traveling semidiurnal internal tidal waves encounter. Following numerical simulations by *Korobov and Lamb* [2008], our observations confirm that PSI of the IT can be effectively enhanced in the reflection of ITs from the boundary. There, PSI is likely to be a strong, rapid nonlinear instability, rather than the classic weakly nonlinear wave-wave interaction. In addition to subharmonics, a mean flow, along with higher harmonics, is also generated by the interaction between two vertically oppositely traveling waves. Although the presence of internal tidal beams cannot be confirmed from our current data, the theoretical framework by *Tabaei et al.* [2005] may be used to explain generation of higher harmonics and the mean flow. Their theory, however, does not explain the occurrence of strongly rapid PSI in the reflection area of internal waves. For the mean flow, some other interesting phenomena seen in the laboratory experiments by *Mercier et al.* [2012] cannot be identified due to the limitation of our data. For these, further study and more data are necessary.

[17] **Acknowledgments.** We thank Dongxiao Wang for the data used in this study. This work is supported by projects 2013CB430303, XDA10010104, U1033002 and 41276021.

[18] Meric A. Srokosz thanks two anonymous reviewers for their assistance in evaluating this paper.

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